

C5.2 Elasticity and Plasticity

Lecture 6 — Elastic waves

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Elastic waves

- ▶ Recall the **unsteady** Navier equation (no body force)

$$\rho \frac{\partial^2 \mathbf{u}}{\partial t^2} = (\lambda + \mu) \text{grad div } \mathbf{u} + \mu \nabla^2 \mathbf{u}$$

- ▶ Seek **travelling harmonic wave** solutions of the form

$$\mathbf{u}(\mathbf{x}, t) = \mathbf{a} \exp \left[i(\mathbf{k} \cdot \mathbf{x} - \omega t) \right] \quad (\text{NB real part assumed})$$

$$\mathbf{a} = (\text{complex}) \text{ amplitude} \quad |\mathbf{k}| = \text{wavenumber} = \frac{2\pi}{\text{wavelength}}$$

$$\omega = \text{frequency} \quad \frac{\mathbf{k}}{|\mathbf{k}|} = \text{wave propagation direction}$$

$$\mathbf{k} = \text{wave-vector} \quad c = \frac{\omega}{|\mathbf{k}|} = \text{wave-speed}$$

Elastic waves

- ▶ Plug $\mathbf{u}(\mathbf{x}, t) = \mathbf{a}e^{i(\mathbf{k}\cdot\mathbf{x}-\omega t)}$ into unsteady Navier equation.
- ▶ There are two distinct types of solution (Problem sheet 2)

(i) P-waves

“Primary” or “Pressure” waves

- ▶ **Longitudinal** waves with $\mathbf{a} \times \mathbf{k} = \mathbf{0}$, i.e.

$$\mathbf{u} = A\mathbf{k}e^{i(\mathbf{k}\cdot\mathbf{x}-\omega t)}$$

- ▶ **P-wave speed**

$$\frac{\omega}{|\mathbf{k}|} = c_p = \sqrt{\frac{\lambda + 2\mu}{\rho}}$$

(ii) S-waves

“Secondary” or “Shear” waves

- ▶ **Transverse** waves with $\mathbf{a} \cdot \mathbf{k} = 0$, i.e.

$$\mathbf{u} = (\mathbf{B} \times \mathbf{k})e^{i(\mathbf{k}\cdot\mathbf{x}-\omega t)}$$

- ▶ **S-wave speed**

$$\frac{\omega}{|\mathbf{k}|} = c_s = \sqrt{\frac{\mu}{\rho}}$$

P-waves and S-waves

P-wave speed

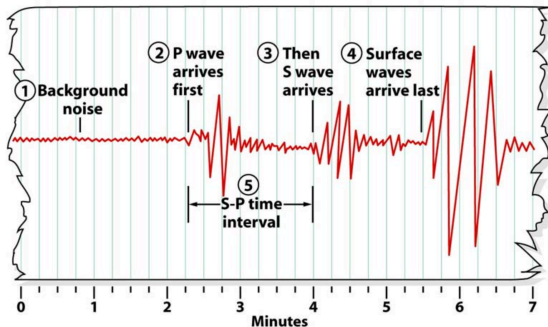
$$c_p = \sqrt{\frac{\lambda + 2\mu}{\rho}}$$

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S-wave speed

$$c_s = \sqrt{\frac{\mu}{\rho}}$$

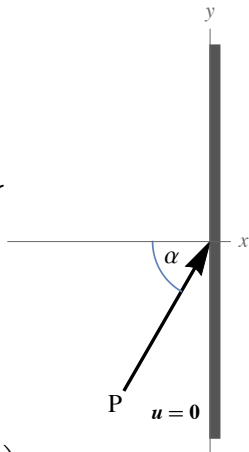
- ① The earthquake happens at time 0.
- ② The first P waves arrive a little over 2 minutes later.
- ③ The first S waves arrive 4 minutes later.



- ④ The surface waves, which travel the long way around Earth's surface, arrive last.
- ⑤ The S-P interval, here slightly less than 2 minutes, tells the seismologist how far away the earthquake was.

Wave reflection

- ▶ Solid occupies half-space $x < 0$ with a straight **rigid** boundary at $x = 0$.
- ▶ P-wave travels towards the boundary making **incident angle** α .
- ▶ NB can generalise, e.g. incident S-wave or stress-free instead of rigid boundary.



Incident wave-field

$$\mathbf{u}_{\text{inc}} = \begin{pmatrix} \cos \alpha \\ \sin \alpha \end{pmatrix} e^{ik_p(x \cos \alpha + y \sin \alpha) - i\omega t}$$

$$\begin{pmatrix} \cos \alpha \\ \sin \alpha \end{pmatrix} = \text{direction vector}$$

$$k_p \begin{pmatrix} \cos \alpha \\ \sin \alpha \end{pmatrix} = \text{wave-vector}$$

amplitude = 1 (wlog)

$$k_p = \frac{\omega}{c_p} = \frac{\omega}{\sqrt{(\lambda + 2\mu)/\rho}}$$

Wave reflection

- ▶ Now include reflected wave-field \mathbf{u}_{ref} s. t.

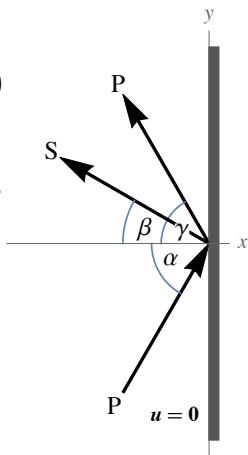
$$\mathbf{u} = \mathbf{u}_{\text{inc}} + \mathbf{u}_{\text{ref}} = \mathbf{0} \text{ at } x = 0 \quad (\star)$$

- ▶ \mathbf{u}_{ref} must consist of both P-wave and S-wave to satisfy both components of (\star) .
- ▶ Consider P- and S- waves making reflected angle γ and β , respectively.

$$\mathbf{u}_{\text{ref}} = r_p \begin{pmatrix} -\cos \gamma \\ \sin \gamma \end{pmatrix} e^{ik_p(-x \cos \gamma + y \sin \gamma) - i\omega t} \\ + r_s \begin{pmatrix} \sin \beta \\ \cos \beta \end{pmatrix} e^{ik_s(-x \cos \beta + y \sin \beta) - i\omega t}$$

$$\text{wave-vectors} = k_p \begin{pmatrix} -\cos \gamma \\ \sin \gamma \end{pmatrix} \text{ and } k_s \begin{pmatrix} -\cos \beta \\ \sin \beta \end{pmatrix}$$

$$k_s = \frac{\omega}{c_s} = \frac{\omega}{\sqrt{\mu/\rho}} \quad r_p, r_s = \text{reflection coefficients}$$



Wave reflection

Now impose $\mathbf{u} = \mathbf{u}_{\text{inc}} + \mathbf{u}_{\text{ref}} = \mathbf{0}$ at $x = 0$ (NB $e^{i\omega t}$ is a factor)

$$\begin{pmatrix} \cos \alpha \\ \sin \alpha \end{pmatrix} e^{ik_p y \sin \alpha} + r_p \begin{pmatrix} -\cos \gamma \\ \sin \gamma \end{pmatrix} e^{ik_p y \sin \gamma} + r_s \begin{pmatrix} \sin \beta \\ \cos \beta \end{pmatrix} e^{ik_s y \sin \beta} = \mathbf{0}$$

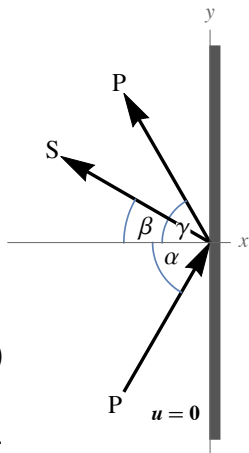
- ▶ Must hold $\forall y$ so **exponents must agree**
- ▶ Therefore P-wave reflection is **specular**

$$\gamma = \alpha$$

- ▶ S-wave reflection satisfies **Snell's law**

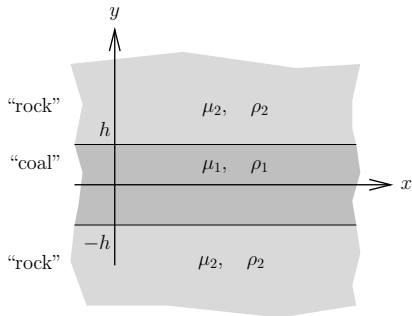
$$\frac{\sin \beta}{c_s} = \frac{\sin \alpha}{c_p} \quad (c_s < c_p)$$

- ▶ Then solve 2×2 linear system for (r_p, r_s)
- ▶ **Mode conversion** — boundaries convert pure P- or S-wave to combination of both.



Love waves

- ▶ Example of wave propagation in a **layered** elastic medium. . .
- ▶ e.g. an underground coal seam.



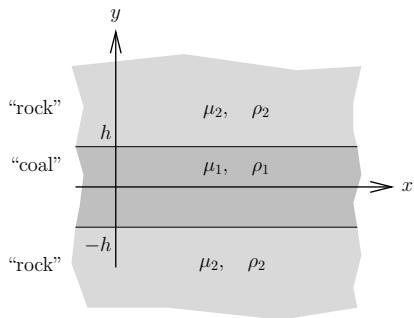
- ▶ Consider **antiplane** displacement

$$\mathbf{u} = \begin{pmatrix} 0 \\ 0 \\ w(x, y, t) \end{pmatrix}$$

- ▶ Unsteady Navier equation $\Rightarrow w$ satisfies the **wave equation**

$$\frac{\rho}{\mu} \frac{\partial^2 w}{\partial t^2} = \boxed{\frac{1}{c_s^2} \frac{\partial^2 w}{\partial t^2} = \nabla^2 w} = \frac{\partial^2 w}{\partial x^2} + \frac{\partial^2 w}{\partial y^2}$$

Love waves



$$\frac{1}{c_{s1}^2} \frac{\partial^2 w}{\partial t^2} = \nabla^2 w, \quad |y| < h$$

$$\frac{1}{c_{s2}^2} \frac{\partial^2 w}{\partial t^2} = \nabla^2 w, \quad |y| > h$$

$$c_{si} = \sqrt{\frac{\mu_i}{\rho_i}}$$

- ▶ Continuity of displacement and stress at boundaries:

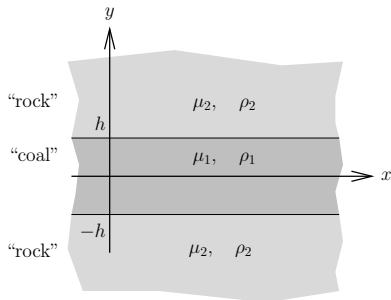
$$w_1 = w_2, \quad \mu_1 \frac{\partial w_1}{\partial y} = \mu_2 \frac{\partial w_2}{\partial y} \quad \text{at } y = \pm h.$$

Love waves

Seek travelling-wave solutions

$$w_i(x, y, t) = f_i(y) e^{i(kx - \omega t)}$$

Suppose displacement is **sinusoidal** in coal seam and **decays exponentially** at infinity:



$$f_1(y) = A_1 \cos(my) + B_1 \sin(my) \quad -h < y < h$$

$$f_2(y) = A_2 e^{-\ell y} \quad y > h$$

$$f_2(y) = B_2 e^{\ell y} \quad y < -h$$

The seam acts as a **waveguide**, propagating waves in the x -direction without energy radiating to ∞ .

Love waves

- ▶ Substitute assumed solution into the wave equation to get...

$$\omega^2 = c_{s1}^2 (k^2 + m^2) = c_{s2}^2 (k^2 - \ell^2)$$

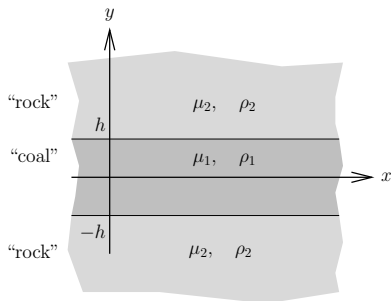
- ▶ Love waves travel at a wave-speed c_L satisfying...

$$\frac{\omega^2}{k^2} = c_L^2 = c_{s1}^2 \left(1 + \frac{m^2}{k^2} \right) = c_{s2}^2 \left(1 - \frac{\ell^2}{k^2} \right)$$

- ▶ Such waves exist provided $c_{s1} < c_L < c_{s2}$ i.e. coal must be **slower** than rock.

Love waves

- ▶ Apply boundary conditions $[f]_{-}^{+} = [\mu f']_{-}^{+} = 0$ at $y = \pm h$
- ▶ Solutions for $f(y)$ can be either **even** or **odd**...
- ▶ **Even** solution: $B_1 = 0, B_2 = A_2$
- ▶ **Odd** solution: $A_1 = 0, B_2 = -A_2$



$$f_1(y) = A_1 \cos(my) \quad |y| < h$$

$$f_2(y) = A_2 e^{-\ell y} \quad y > h$$

$$f_2(y) = A_2 e^{\ell y} \quad y < -h$$

Even modes:

$$\mu_1 m \tan mh = \mu_2 \ell$$

Odd modes:

$$\mu_1 m \cot mh = -\mu_2 \ell$$

Love waves

- ▶ Transcendental equation relating m and ℓ :

$$\frac{\mu_2 \ell}{\mu_1 m} = \begin{cases} \tan(mh) & \text{even modes} \\ -\cot(mh) & \text{odd modes} \end{cases}$$

- ▶ Simplest case: $\mu_2/\mu_1 \rightarrow \infty$ i.e. rock **much harder** than coal.
- ▶ Then $mh = (n+1)\pi/2$ where $n \in \mathbb{Z}_{\geq 0}$;
 n is **odd** for odd modes and **even** for even modes.
- ▶ Then $\frac{\omega^2}{k^2} = c_L^2 = c_{s1}^2 \left(1 + \frac{(n+1)^2 \pi^2}{4h^2 k^2} \right)$
- ▶ Love waves are **dispersive** — c_L varies with k .
- ▶ Note **cut-off frequency** $\omega > \omega_c = \frac{\pi c_{s1}}{2h}$
waves cannot propagate without attenuation if $\omega < \omega_c$